Long-term Atmospheric Mercury Wet Deposition at Underhill, Vermont

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Abstract. Section 112(m) of the 1990 Clean Air Act Amendments, referred to as the Great Waters Program, mandated an assessment of atmospheric deposition of hazardous air pollutants (HAPs) to Lake Champlain. Mercury (Hg) was listed as a priority HAP and has continued to be a high priority for a number of national and international programs. An assessment of the magnitude and seasonal variation of atmospheric Hg levels and deposition in the Lake Champlain basin was initiated in December 1992 which included event precipitation collection, as well as collection of vapor and particle phase Hg in ambient air. Sampling was performed at the Proctor Maple Research Center in Underhill Center, VT. The range in the annual volume-weighted mean concentration for Hg in precipitation was 7.8–10.5 ng/l for the 11-year sampling period and the average amount of Hg deposited with each precipitation event was $0.10 \ \mu g/m^2$. The average amount of Hg deposited through precipitation each year from 1993 to 2003 was 9.7 $\mu g/m^2/yr$. A seasonal pattern for Hg in precipitation is clearly evident, with increased Hg concentrations and deposition observed during spring and summer months. While a clear trend in the 11-year event deposition record at Underhill was not observed, a significant decrease in the event max-to-monthly ratio was observed suggesting that a major source influence was controlled over time. Discrete precipitation events were responsible for significant fractions of the monthly and annual loading of Hg to the forested ecosystem in Vermont. Monthly-averaged temperatures were found to be moderately correlated with monthly volumeweighted mean Hg concentrations ($r^2 = 0.61$) and Hg deposition ($r^2 = 0.67$) recorded at the Vermont site. Meteorological analysis indicated the highest levels of Hg in precipitation were associated with regional transport from the west, southwest, and south during the warmer months.

Keywords: trends; seasonal patterns; speciation; photochemistry

Introduction

Mercury (Hg) is a toxic bioaccumulative substance found in aquatic ecosystems. Atmospheric deposition is widely recognized as an important link in the cycling of Hg in the environment (Lindqvist and Rhode, 1985; Mierle, 1990; Fitzgerald et al., 1991; Iverfeldt, 1991). Lake Champlain, located in the

*To whom correspondence should be addressed: Tel.: +1-734-936-1836 E-mail: jkeeler@umich.edu northeastern United States on the border between northern New York State, Vermont, and Quebec, has been designated as one of the "Great Waters" by the 1990 Clean Air Act Amendments (CAA). The Great Waters Program was established in Section 112(m) of the CAA to identify and assess atmospheric deposition of hazardous air pollutants (HAPs) to the Great Lakes, Chesapeake Bay, Lake Champlain and coastal waters. Mercury (Hg) is one of the 189 HAPs identified in this section of the CAA. Requirements of the Great Waters program

include: (a) establishing atmospheric deposition stations to monitor deposition of HAPs within the Lake Champlain watershed, (b) determining the role of atmospheric deposition in the pollutant loading for the lake, and (c) investigating the sources of air pollutants deposited in the watershed.

Mercury concentrations exceeding fish consumption advisory limits have been documented for certain sport fish, and elevated levels of Hg and PCBs have been found in the sediments at many sites within Lake Champlain including those in the deep lake (McIntosh, 1994). This report identified several lake-wide issues for future investigation, including the critical need to better understand how contaminants like Hg and PCBs are entering Lake Champlain. More recently the New England Governors and Canadian Premiers (2003) have identified that Hg releases to the environment with "resulting public health and environmental impacts" need to be addressed.

Atmospheric deposition has been implicated as a primary source of Hg to remote lakes in the Great Lakes region (Fitzgerald et al., 1991; Swain et al., 1992). Atmospheric Hg deposition has also been shown to play a significant role in Hg loading to fresh water ecosystems and watersheds in Vermont and New England (Scherbatskoy et al., 1994, 1998; Kamman and Engstrom, 2002). While atmospheric deposition of acidic species and transport of these pollutants from the Midwest to New England has been documented (Butler et al., 1991), relatively few studies have documented the transport and deposition of Hg to northern New England (Scherbatskoy et al., 1994; Burke et al., 1995; Macolm et al., 2003). These Hg studies were all part of a coordinated long-term study of Hg cycling in the Lake Champlain watershed. Here we report on the only long-term event precipitation Hg data collected in the New England region, and the factors that influenced the atmospheric concentrations and deposition of this persistent, bioaccumulative contaminant.

Collection and analysis methods

Clean technique and acid cleaning procedure

Precipitation samples were processed and analyzed in a Class 100 clean room at the University of Michigan to avoid potential contamination. Clean room suits and particle-free gloves were worn at all times during preparation and analysis of samples. All field and analytical supplies used in the collection and analysis of Hg and trace element samples were prepared using an 11-day acidcleaning procedure (Landis and Keeler, 1997) to ensure that all materials that potentially came in contact with a sample were essentially Hg-free. Standard operating procedures included bottle blank determinations, which ensured that sampling bottles were Hg-free before they were used in the field. Furthermore, the entire precipitation sampling train, for both Hg and trace elements, was replaced with a freshly acid-cleaned set after each precipitation event.

Wet deposition sampling

The sampling site is located on the west slope of Mount Mansfield at the Proctor Maple Research Center (PMRC) at 400 m elevation. The PMRC is located in the Lake Champlain basin approximately 25 km east of the lake. No major urban or industrial areas exist within 200 km of the site. Wet-only event precipitation samples were collected using a modified MIC-B (MIC, Thornhill, Ontario) automatic precipitation collector (Landis and Keeler, 1997). During the first 2 years of the study samples were collected into a 101 borosilicate glass (BSG) vessel using the MIC-B with a Teflon-coated funnel which had an inlet area of 0.2 m² (Hoyer et al., 1995). Samples were transferred from the 101 BSG vessel into a 11 BSG bottle immediately after collection. This approach was found to be difficult to implement on a large scale and the glass bottles were not as desirable due to breakage during shipment. Glass bottles were also not considered the best material for the collection of the other trace elements. In early summer of 1994 the system described in Landis and Keeler (1997) replaced the original one and remains in use. Enrichment and adsorption of trace metals in precipitation was minimized by the use of separate Hg and trace element sampling trains composed of different materials (Ross, 1986, 1990). The Hg sampling train was comprised of a borosilicate glass funnel (collection area of $191 \pm 9 \text{ cm}^2$), and a Teflon adapter, which incorporated a glass vapor lock, attached to a 1 l Teflon sample bottle. The trace element sampling train was comprised of a polypropylene funnel (collection area of $167 \pm 7 \text{ cm}^2$) and a polypropylene adapter attached to a 1-l polypropylene sample bottle. The collection of Hg and trace elements directly into separate sample bottles was also important due to the absorptive behavior of trace metals in precipitation to the walls of the bottles (Church et al., 1984). While samples for both trace elements and Hg were collected for each event, only the long-term Hg record is reported in this manuscript.

Mercury samples were oxidized with concentrated BrCl to a 1% solution (v/v) and were stored in a dark cold room before being analyzed. Post oxidation of the Hg samples and acidification of the trace element samples to a pH <1.5 prohibits absorptive loss to the container walls (Subramanian et al., 1978; Scudlark et al., 1991). The sampling trains were supported in the UM Modified MIC-B collector by a UM custom-built acrylic insert. The volume of each precipitation sample was determined gravimetrically. Precipitation events ≥ 0.1 cm in depth were efficiently collected by the MIC-B. Precipitation only overflowed the collection bottle for relatively large events > 5.6 cm. The funnels had high vertical sides to aid in efficient collection and to prevent loss due to "bounce off" during intense rain. In addition, the sampling trains were designed to be light so they could be easily removed, disassembled and acid cleaned. During the coldest months of the year (November-April) a small ceramic space heater (750 W) was installed inside the base of the collector to maintain a cabinet temperature of ~10 °C to melt snow that collected in the funnels and prevent collected samples from freezing. A Belfort precipitation gauge provided a continuous record of the precipitation received at the Underhill site. The precipitation depth for the first 2-years of the study was determined directly from the Belfort charts, and in subsequent years the volume of precipitation collected in the glass funnel and the area of the collection funnel was used to calculate the precipitation depth. The Belfort precipitation gauge also provided the actual rain amount in the rare event that the 1-l sample bottles overflowed due to extreme rain events.

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The total Hg concentration in precipitation was determined by purging Hg from solution onto a gold-coated bead trap using a Hg-free nitrogen stream after pre-reduction of BrCl with NH₂OH, and reduction of divalent Hg by SnCl₂ to elemental gaseous Hg (Hg⁰). Total Hg was quantified by a dual amalgamation technique followed by cold-vapor atomic fluorescence spectrometry (CVAFS) (Landis and Keeler, 1997). Co-located total Hg samples were collected using identical samplers and protocols as part of the Lake Michigan Mass Balance Study (LMMBS). The LMMBS ran concurrently for 18-months with the long-term monitoring conducted at the Underhill site. An absolute mean difference of 8.1% was calculated between the co-located samples (Landis and Keeler, 1997). Analytical precision of laboratory replicates during this study was 95.5%.

Meteorological trajectory analysis

The transport of air to the Underhill site was estimated using the Hybrid Single-Particle Lagrangian Integrated Trajectory (HYSPLIT) Model Version 4.6 (Draxler and Hess, 1997). HYSPLIT 72-h back trajectories were calculated using input data from The National Weather Service's National Center for Environmental Prediction (NCEP) Nested Grid Model (NGM) for 1993-1996 and the Eta Data Assimilation System (EDAS) for 1997–2003. The NGM and EDAS data were obtained from a standard data archive maintained by the National Oceanic and Atmospheric Administration's Air Resources Laboratory (NOAA-ARL). NOAA-ARL routinely archives NGM and EDAS data in a format that can be read by the HYSPLIT Modeling System.

The hour of maximum precipitation measured from each event was used as the start time for the trajectory calculation. The starting height for each trajectory was calculated as one-half the mixing height determined from the upper-air soundings. Surface and upper air meteorological maps obtained from the National Weather Service (NWS) were used to explore the validity of the calculated trajectories and to better understand the type of precipitation (convective, frontal, etc.) that was received.

Results

Event wet deposition

The University of Michigan Air Quality Laboratory (UMAQL) has operated the only long-term event precipitation sites in the United States, which began in the early 1990s. Precipitation collection at the PMRC site began in December 1992 and has continued without interruption to the present. The concentration of Hg measured in event samples (shown in Fig. 1) includes data for the 11-year period from 1/1/1993 to 12/31/2003. Event precipitation concentrations varied by over two orders of magnitude with a range in individual measured event concentrations of $\sim 0.9-90$ ng/l. The volume-weighted mean (VWM) concentration for the data shown was 8.9 ng/l (N = 1111). The average event, over the 11-years, had a precipitation depth of 1.1 cm and a maximum depth of 6.2 cm.

While Hg concentrations were influenced by the volume of precipitation collected, the precipitation depth could only explain approximately 20% of the variance in the Hg concentration. Other factors such as the season, type of precipitation, and meteorological transport history more strongly influenced the Hg concentration observed at Underhill. The occurrence of precipitation events with relatively low concentrations (<2.0 ng/l) was independent of the precipitation volume collected, suggesting that under certain conditions little Hg was available to be removed or the chemical form of the Hg was not amenable to washout or rainout processes.

The atmospheric loading of Hg to Underhill from 1993 to 2003 is shown in the event deposition plot (Fig. 2). Wet deposition was calculated, on an event basis, by taking the product of the Hg concentration (shown in Fig. 1) and the precipitation depth. Event Hg deposition varied from $\sim 0.002 \ \mu g/m^2$ to 1.8 $\mu g/m^2$ with the average event wet deposition for the 11-year time period being 0.10 $\mu g/m^2$.

The importance of individual precipitation events to the monthly and yearly Hg deposition was clearly seen in the 11-year record. The single largest deposition event, which delivered 1.8 μ g/m², was recorded at Underhill on 21 October 2003. This unusual event was the product of a 5.1 cm rain, associated with a low pressure system sitting over northern New York State, with an above average Hg concentration of 35.1 ng/l. High pressure centered over central Pennsylvania on 20 October 2003 resulted in stagnation and low-mixing heights over western New York, Pennsylvania, Ohio, and southern Ontario. Air mass transport into northern New York and Vermont, ahead of the warm front attached to



Figure 1. Event Hg concentrations measured at Underhill, Vermont from 1993-2003.



Figure 2. Event Hg deposition measured at Underhill, VT from 1993-2003.

the low pressure system, was from the westsouthwest from the areas of stagnation on the previous day. Interestingly, a large precipitation event with 5.3 cm of rainfall was also recorded only 5 days earlier, but had a below average concentration of 3.8 ng/l, resulting in a Hg deposition of only 0.2 μ g/m². These single large deposition events (>0.4 μ g/m²) can account for 20–60% of the monthly deposition, 5–17% of the total annual deposition recorded at this site, and 13% of the total Hg deposition recorded at the Underhill site over the 11-year period.

A significant trend was not evident in the maximum event Hg deposition, or the arithmetic mean deposition, over the 11 years of observation reported here. Interestingly, there was a statistically significant decrease (19%) in the maximum event deposition to monthly mean deposition ratio over the 11 years ($r^2 = 0.37$). A similar decrease in the max-to-monthly mean Hg concentration ratio was also observed ($r^2 = 0.28$) perhaps suggesting that a source of the elevated event Hg concentrations and deposition has been controlled. A meaningful examination of the chemical and meteorological parameters that control the removal of Hg species from the atmosphere requires the precipitation samples to be collected in discrete events (Sisterson et al., 1985; Ross, 1990; Dvonch et al., 1999).

Seasonal variability

The key to understanding the spatial and temporal patterns in atmospheric deposition lies in our ability to explain the seasonality in Hg deposition. The event Hg deposition data, shown in Fig. 2, displays clear seasonal trends. Higher Hg deposition, typically observed during the warmer months, was likely the result of a mix of meteorological, source emission, and atmospheric chemistry influences. For example, it is widely known that the concentrations of oxidants such as ozone, OH radicals, and acids that oxidize gaseous Hg^0 to gaseous Hg^{2+} are higher during warmer months and would lead to elevated concentrations of Hg²⁺ (Schroeder and Menthe, 1998; Lin and Pekonen, 1999). Scavenging of soluble gaseous Hg^{2+} has also been pointed to as more efficient in summertime precipitation events than winter due to differences in the cloud microphysical processing between rain and frozen precipitation (Hoyer et al., 1995; Keeler et al., 1997). The warmer months also had a higher baseline (minimum values) in the event Hg deposition than the colder months as seen in Fig. 2. The lowest concentrations and total wet deposition were seen in winter months, and highest values appeared in the spring and summer months. This may, in part, indicate the role of convective transport of Hg to

this region, and may also be reflective of the types of clouds and storms that bring precipitation to northern Vermont during the summer relative to the winter. The 11-year plot for total monthly Hg wet deposition recorded at the PMRC site is shown in Fig. 3. The monthly volume-weighted concentration plot, not shown, looked similar to the deposition plot with similar monthly trends and year-to-year differences in the monthly concentrations.

While little year-to-year difference was observed in the minimum monthly Hg wet deposition, the maximum monthly total varied greatly from year to year. Calendar year 1998 was a particularly significant year for Hg wet deposition. This was largely due to three well-above average deposition months (June–August) during the spring–summer period which recorded large individual events (>0.4 μ g/m²) in each of the months. There did not, however, appear to be an overall increase or decrease in the Hg concentrations over time. This will be examined more closely in the subsequent discussion.

The seasonal variability in Hg concentrations and Hg deposition has been reported in previous studies in North America (Burke et al., 1995; Hoyer et al., 1995; Landis and Keeler, 1997) and Europe (Iverfeldt, 1991). The larger data set from Underhill provides a clearer picture of how the Hg levels fluctuate seasonally over the 11 annual cycles. The seasonal pattern in the atmospheric Hg, with highest precipitation concentrations and wet deposition typically seen in the summer and lowest concentrations and wet deposition in the winter, was believed partly to be the result of increased convection during the warmer summer months which can increase the ability of the air to transport Hg over longer distances, and leads to greater precipitation amounts that remove Hg from the atmosphere. This may also indicate the role of precipitation type in the amount of Hg wet deposition, as rain may have a greater capacity to scavenge and hold different forms of Hg than snow.

To examine the relationship between concentration and season, we calculated the frequency distribution for the 11-year event precipitation record. Fig. 4 shows the frequency distribution for Hg concentrations in each season over the 11-year period. For each sample collected, the number of samples that fell into each concentration bin, 0-5, 5-10, 10-15, 15-20, etc. was determined and then subdivided into the time of year that the sample was collected. Event samples that were collected between 1993-2003 were most often (>31%) in the 5–10 ng/l concentration range and were just as likely to be collected in any of the four seasons. The next most frequent concentrations observed were in the 0-5 ng/l range with the majority of these samples from the autumn and



Figure 3. Monthly Total Hg wet deposition at Underhill, VT.

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Figure 4. Frequency distribution of seasonal Hg concentration measured in precipitation at Underhill, VT from 1993–2003 (N = 1111).

winter seasons. The highest concentrations of Hg were more common during the spring and summer seasons rather than in winter or autumn. Samples collected during the summer were more likely to have concentrations in the 10–15 ng/l range.

The frequency distribution for Hg measured in event wet deposition samples for each season is shown in Fig. 5. Lower Hg wet deposition amounts were more frequent during the autumn and winter, and higher Hg values were typically present during the spring and summer. The majority (>43%) of all event samples collected resulted in Hg deposition in the range 0–0.05 μ g/m² with another 25% of the events having values in the 0.05–0.1 μ g/m² range. Less than 5% (N = 39) of the event samples resulted in Hg deposition > 0.4 μ g/m², however, these events contributed to 19% of the total deposition recorded at the



Figure 5. Frequency distribution of seasonal Hg wet deposition measured in precipitation at Underhill, VT from 1993–2003 (N = 1111).

Underhill site over the 11-year period. Each of these events had both above average Hg concentrations as well as above average precipitation depths. All of the 39 precipitation events that resulted in Hg deposition values above $0.4 \ \mu g/m^2$ were in the form of rain. Of these events, 24 fell during the summer months, 11 during the spring and 4 during the autumn. These high levels of Hg were much less frequent, but point to the importance of atmospheric chemistry, cloud microphysics, and transport during the spring and summer months. It was the frequency of these elevated Hg deposition events that resulted in above average annual deposition amounts. Thus, understanding these events may lead to a better understanding of the year-to-year patterns in deposition.

Annual variability

Considerable sample to sample variability observed in the daily event data can be explained largely by variations in the meteorological processes and source-receptor relationships that influence the concentrations and deposition of Hg. Studies assessing changes in atmospheric deposition over time scales longer than seasons suggest that short-term meteorological variability is less significant than longer-term climatic variability (Bradley, 1986). While long-term trends in Hg deposition may be more strongly influenced by climatic fluctuations, the importance of large events appears to be quite significant in controlling the annual levels of atmospheric Hg deposition at the northern Vermont site.

The annual volume-weighted mean (VWM) Hg concentration and total Hg wet deposition collected at the Underhill site are summarized in Table 1. The number of events collected (>0.1 cm) each year of sampling at Underhill ranged from 71 to 124 over the 11-years of observation (See Table 1). The maximum values in any column are shown in bold for emphasis. While the number of events per year varied, on average, by only 15% from the 11-year mean, the type and intensity of the events were quite different, resulting in precipitation depths that varied by more than a factor of two. The number of small events that were not collected for analysis also varied considerably from year to year with the average number equal to 13, and a range of 4-27 small events per year. The years 2000, 2002, and 2003 all recorded more than 25 precipitation events < 0.1 cm in precipitation depth.

The data in Table 1 indicate that 1998 was a particularly significant year for precipitation, totaling 141 cm. The year 2003 also shows a high precipitation total of 133 cm. Although these two years had similar precipitation amounts, there were many more events in 1998 than in 2003. The year 1998 recorded 123 events while 2003 saw only 90, indicating a greater intensity of the events in 2003. In addition, 1993 and 1998 both received 124

Table 1. Annual volume-weighted mean (VWM) Hg concentration and total Hg wet deposition from event samples collected at Underhill, VT from 1993–2003

| Year | Precipitation depth (cm) | Total Hg wet deposition $(\mu g/m^2)$ | Average Hg wet deposition $(\mu g/m^2)$ | VWM Hg concentration (ng/l) | # events collected | Maximum concentration (ng/l) | Maximum deposition (µg/m ²) |
|------|-----------------------------|---------------------------------------|-----------------------------------------------|-----------------------------------|--------------------|---------------------------------|--------------------------------------------|
| 1993 | 103.2 | 10.8 | 0.09 | 10.5 | 124 | 51.4 | 0.50 |
| 1994 | 97.9 | 8.6 | 0.08 | 8.8 | 115 | 43.7 | 0.55 |
| 1995 | 105.9 | 8.7 | 0.09 | 8.2 | 99 | 59.1 | 0.66 |
| 1996 | 102.7 | 8.4 | 0.08 | 8.2 | 102 | 89.8 | 0.48 |
| 1997 | 97.2 | 9.3 | 0.10 | 9.6 | 98 | 53.1 | 0.74 |
| 1998 | 141.1 | 13.3 | 0.11 | 9.5 | 124 | 53.9 | 0.65 |
| 1999 | 109.1 | 8.6 | 0.08 | 7.8 | 107 | 49.0 | 0.41 |
| 2000 | 105.5 | 9.0 | 0.10 | 8.5 | 95 | 51.1 | 0.45 |
| 2001 | 85.0 | 8.6 | 0.12 | 10.1 | 71 | 38.0 | 0.61 |
| 2002 | 118.9 | 10.2 | 0.12 | 8.6 | 85 | 40.9 | 0.60 |
| 2003 | 133.1 | 10.9 | 0.12 | 8.2 | 91 | 45.7 | 1.79 |
| MEAN | 109.1 | 9.7 | 0.10 | 8.9 | 101 | 52.3 | 0.68 |

events, the highest of any year. However, 1998 recorded a much higher precipitation total, implying that the 1998 events were more intense that the 1993 events. In contrast, 2001 had both lowest yearly precipitation depth (85 cm) and the smallest number of events (71).

The volume-weighted mean Hg concentration varied over each year as well, with a maximum value of 10.0 ng/l calculated for 2001 and a minimum of 7.8 ng/l in 1999. The volume-weighted mean for all years was 8.7 ng/l and the average total Hg deposition for the eleven-year period was 9.7 μ g/m². It is important to note the high volumeweighted mean concentrations in 1993, 1997, 1998, and 2001. While 1998 received the highest precipitation of any year over the 11-years, 1997 and 2001 received less than average (=109.1 cm) precipitation amounts. During 1998, 13.3 μ g/m² of Hg deposition was recorded, and this was also the second highest deposition year for sulfate, nitrate, chloride, and other major ions in precipitation reported by the National Acid Deposition Program (NADP). However, unlike the apparent downward trend in sulfate deposition at Underhill, the annual Hg levels observed at Underhill show variation from year to year without a clear increasing or decreasing trend. To further understand the possible causes for the seasonal and annual variations, the Hg data was compared with other long-term Hg records, as well as with the on-site meteorological data and the upwind history of each event determined by back trajectories.

Comparison with other long-term event records

Recognizing that long-term precipitation records are essential for establishing trends and understanding the impacts of changes in Hg emissions, the UMAQL also collected a decade of event precipitation data at three sites in Michigan (Dexter, Pellston, and Eagle Harbor) for the period 1994–2003. Over this 10-year deposition record, a clear decreasing gradient from south to north was observed in Michigan. The 10-year Hg deposition total recorded at Underhill from 1994– 2003 compared closely with the Pellston site Hg deposition, as the two sites collected 96 versus 92 μ g/m², respectively. The deposition at these two sites was greater than the deposition observed at Eagle Harbor, on the shores of Lake Superior, which was 58 μ g/m², but lower than the deposition measured at Dexter, MI of $120 \ \mu\text{g/m}^2$ for the 10 years. While the year-to-year variability in the deposition was, on average, 18% at the Michigan sites, the 10-year variability observed at Underhill was only 12%. Interestingly, there was not an obvious trend in the Hg deposition at three Michigan sites (Keeler and Dvonch, 2004) or at the Vermont site over the 10 years from 1994-2003. While there have been attempts to control emissions within the region and nationally over the past decade (Butler et al., 2001), the MI data illustrates the consistent long-term impact that sources in the southern part of the Great Lakes region have had on Hg deposition across the Great Lakes Basin.

Temperature dependence

The monthly-averaged Hg data was examined in relationship to temperature data obtained from the meteorological tower at the Proctor Maple Research Center. Temperatures for each day with a precipitation event were averaged monthly for comparison with the monthly VWM Hg concentration and monthly total wet deposition data. Figure 6 displays the average temperature (line) and average Hg deposition (bars) for each month over the 11-year period. This figure suggests that the monthly variation in the temperature strongly influences the processes that control the deposition of Hg from the atmosphere. As previously mentioned, the transport of Hg may be higher with warmer temperatures due to increased convection and mixing. The plot shows that higher levels of Hg were deposited during the warmer, more photochemically active summer months, and lower amounts of Hg were deposited in the colder months. A significant relationship between monthly temperature and Hg wet deposition was observed with an r^2 value of 0.67.

In addition, the on-site meteorological data, including the hourly temperature measurements, were averaged for each day of precipitation. While the monthly averaged temperature and monthly Hg deposition were moderately correlated, this relationship was not as clear when we matched the daily averaged temperature with the resulting daily-event Hg deposition. This implies that either



Figure 6. Monthly total Hg wet deposition and monthly average temperatures (line).

that the meteorological parameters and atmospheric chemistry that varied seasonally largely control the deposition of Hg, or that the upwind temperature and conditions are more important than the temperatures that were recorded on the days that precipitation was collected.

Back trajectory analysis

The importance of the large deposition events on the long-term loading of Hg to the northern Vermont site was discussed earlier in this paper. To further understand the transport of these high levels of Hg, HYSPLIT 72-h back trajectories were calculated for each day of the year. Trajectories associated with precipitation events that resulted in Hg deposition $> 0.4 \ \mu g/m^2$ are shown in Fig. 7. A summary of each of the precipitation events, including the date, Hg concentration, precipitation depth, and total Hg deposition is provided in Table 2.

While the meteorological and source-receptor relationships that were associated with the highest deposition events may be complex, the fact that they largely fell into west-southwest and southerly transport corridors gives us insights into the factors controlling Hg levels and removal from the atmosphere. Poirot et al., 2001 found similar transport and source-receptor relationships using hybrid receptor modeling techniques applied to aerosol data from the PMRC site. They identified the southerly source area as East Coast Oil and the southwesterly source area as a Midwest coal source contribution to the aerosol concentrations measured at Underhill. Each of the events shown had above average Hg concentrations as well as an above average precipitation depth. While air mass transport from the northwest and southeast was also found to bring elevated concentrations of Hg to the Vermont site (Burke et al., 1995), transport from this corridor was generally not associated with the highest Hg deposition events over the 11-year record. Results from hybrid-receptor modeling techniques applied to the long-term event Hg deposition data are forthcoming and should help to identify the dominant source areas contributing to the deposition of Hg in Vermont.

Conclusion

The 11-year record at Underhill, VT is the longest event-based Hg precipitation data set in the United States. The event wet deposition data reveals strong monthly and seasonal trends that imply significant dependence on meteorological conditions throughout the year. The data reveals a temperature dependence that supports the monthly and seasonal variations. The annual analysis indicates little variation in the average transport of Hg to the Underhill region over time, and thus no simple trend

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Figure 7. Back-trajectories associated with precipitation events with deposition events $> 0.4 \ \mu g/m^2$ from Underhill, VT for 1993–2003.

was discerned in the 11-year deposition record. Large deposition events significantly contributed, as much as 17%, to the annual Hg deposition total. It is also apparent that the primary source regions are to the west and southwest of Underhill, VT based on meteorological back-trajectory analysis. Analysis of archived precipitation samples by ICP-MS for trace element concentrations combined with the previously discussed Hg wet deposition data will further our understanding of the atmospheric sources and transport to Northern New England. In addition, a more thorough investigation of the meteorology associated with each event and an indepth examination of the back-trajectories will also be vital to understanding the atmospheric transport and source-receptor relations for Hg. Hybrid receptor models applied to this data will provide a detailed apportionment of the Hg deposition, indicating the most important sources and source regions contributing to the Hg burden at the northern Vermont site.

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| Trajectory | Event | Precipitation | Hg | Hg wet |
|------------|------------|---------------|---------------|---------------|
| ID | start | depth (cm) | concentration | deposition |
| | date | | (ng/l) | $(\mu g/m^2)$ |
| A | 4/25/1993 | 2.1 | 17.7 | 0.4 |
| В | 6/15/1993 | 2.0 | 24.8 | 0.5 |
| С | 7/29/1993 | 1.8 | 21.3 | 0.4 |
| D | 8/24/1993 | 3.6 | 10.4 | 0.4 |
| E | 8/31/1993 | 4.4 | 10.8 | 0.5 |
| F | 9/9/1993 | 2.0 | 22.6 | 0.4 |
| G | 5/31/1994 | 2.4 | 23.0 | 0.6 |
| Η | 6/6/1994 | 3.9 | 12.0 | 0.5 |
| Ι | 5/14/1995 | 1.6 | 30.9 | 0.5 |
| J | 7/1/1995 | 5.6 | 8.7 | 0.5 |
| Κ | 8/3/1995 | 6.0 | 11.0 | 0.7 |
| L | 8/11/1995 | 2.7 | 20.8 | 0.6 |
| Μ | 6/13/1996 | 3.8 | 11.2 | 0.4 |
| Ν | 7/26/1996 | 2.4 | 15.0 | 0.4 |
| 0 | 9/28/1996 | 3.1 | 15.5 | 0.5 |
| Р | 7/9/1997 | 4.9 | 12.3 | 0.6 |
| Q | 7/14/1997 | 3.5 | 21.1 | 0.7 |
| R | 7/15/1997 | 2.0 | 26.2 | 0.5 |
| S | 5/31/1998 | 4.9 | 13.1 | 0.6 |
| Not Shown | 7/16/1998 | 2.6 | 21.0 | 0.5 |
| Т | 8/24/1998 | 2.3 | 17.7 | 0.4 |
| U | 9/15/1998 | 5.2 | 9.4 | 0.5 |
| V | 9/26/1998 | 2.3 | 18.5 | 0.4 |
| W | 9/30/1998 | 2.6 | 13.4 | 0.4 |
| Х | 7/5/1999 | 4.4 | 8.7 | 0.4 |
| Y | 7/9/1999 | 2.3 | 17.8 | 0.4 |
| Ζ | 5/8/2000 | 4.9 | 8.5 | 0.4 |
| AA | 5/18/2000 | 1.7 | 26.8 | 0.4 |
| BB | 6/16/2001 | 3.0 | 13.0 | 0.4 |
| CC | 7/10/2001 | 2.6 | 14.8 | 0.4 |
| DD | 8/31/2001 | 5.0 | 12.2 | 0.6 |
| EE | 6/11/2002 | 4.9 | 8.4 | 0.4 |
| FF | 6/21/2002 | 3.9 | 15.5 | 0.6 |
| GG | 7/5/2002 | 2.6 | 15.2 | 0.4 |
| HH | 7/8/2002 | 4.7 | 12.4 | 0.6 |
| II | 9/14/2002 | 4.0 | 12.3 | 0.5 |
| JJ | 9/22/2002 | 4.4 | 10.6 | 0.5 |
| KK | 8/4/2003 | 4.6 | 14.5 | 0.7 |
| LL | 10/20/2003 | 5.1 | 35.3 | 1.8 |

Table 2. Top deposition events ($>0.4 \mu g/m^2$) at the PMRC site from 1993–2003

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